Accepted refereed manuscript of:

Pickering MD, Ghislandi S, Usai MR, Wilson C, Connelly P, Brothwell DR & Keeley BJ (2018) Signatures of degraded body tissues and environmental conditions in grave soils from a Roman and an Anglo-Scandinavian age burial from Hungate, York. *Journal of Archaeological Science*, 99, pp. 87-98.

DOI: https://doi.org/10.1016/j.jas.2018.08.007

© 2018, Elsevier. Licensed under the Creative Commons Attribution-NonCommercial-NoDerivatives 4.0 International <u>http://creativecommons.org/licenses/by-nc-nd/4.0/</u>

- 1 Signatures of degraded body tissues and environmental conditions in grave soils from a
- 2 Roman and an Anglo-Scandinavian age burial from Hungate, York
- 3 Matthew D. Pickering^a, Sabina Ghislandi^b, Maria Raimonda Usai^{b,c†}, Clare Wilson^d, Peter Connelly^e,
- ⁴ Don R. Brothwell^{b†} and Brendan J. Keely^{a*}
- ⁵ ^a Department of Chemistry, University of York, Heslington, York, YO10 5DD, UK
- ⁶ ^b Department of Archaeology, University of York, The King's Manor, York, YO1 7EP, UK
- 7 ^cDipartimento di Architettura e Design, Wilson, Piazza Pau Salit, Alghero, Italy
- 8 ^dSchool of Biological and Environmental Science, University of Stirling, Stirling. FK9 4LA, UK
- 9 ^eYork Archaeological Trust, 47 Aldwark, York YO1 7BX, UK

10 ORCHID ID

- 11 Keely: <u>0000-0002-8560-1862</u>
- 12 Pickering: <u>0000-0002-6234-2108</u>
- 13 Wilson: <u>0000-0002-0287-8576</u>

14 Highlights

- Signatures of degraded body tissues, gut contents and microbial activity from grave soils.
- Spatial variation in signatures with anatomical location and features of burial environment.
- 17 *n*-Alkanals reveal highly localised regions of anoxia.
- Grave chemistry and preservation of the skeletal remains has been impacted by an adjacent
 cess pit.
- Pedogenic spherulites record broad changes in redox conditions within the Roman grave.
- 21 **Keywords:** Human burials; scientific archaeology; triacylglycerols; *n*-alkanals; bone cholesterol; redox
- 22 conditions; Roman; Anglo-Scandinavian.

⁺ Deceased on 26 September 2016 (DRB) and 13 May 2018 (MRU).

23 Abstract

24 Despite the importance of human burials in archaeological investigations of past peoples and their 25 lives, the soil matrix that accommodates the remains is rarely considered, attention being focused 26 mainly on visible features. The decomposition of a buried corpse and associated organic matter 27 influences both the organic composition and, directly or indirectly, the microstructure of the burial 28 matrix, producing signatures that could be preserved over archaeological timescales. If preserved, 29 such signatures have potential to reveal aspects of the individual's lifestyle and cultural practices as 30 well as providing insights into taphonomic processes. Using organic chemical analysis and soil 31 micromorphology we have identified organic signatures and physical characteristics relating to the 32 presence of the body, and its decomposition in grave soils associated with two human skeletons (one 33 Roman age and one Anglo-Scandinavian age) from Hungate, York, UK. The organic signatures, 34 including contributions from body tissues, gut contents, bone degradation and input from microbiota, 35 exhibit spatial variations with respect to anatomical location and features of the immediate burial 36 In the Roman grave broad changes in redox conditions associated with the environment. 37 decomposition of the corpse and disturbance from the excavation and use of an Anglo-Scandinavian 38 age cess pit that partially cuts the grave were evident. Leachate from the cess pit was shown to exacerbate the degradation of the skeletal remains in the regions closest to it, also degrading and 39 40 depleting spherulites in the soil, through decalcification of the bone and liberation of bone-derived 41 cholesterol into the soil matrix. The findings from this work have implications for future archaeo- and 42 contemporary forensic investigations of buried human remains.

43 1 Introduction

44 The study of human burials in the archaeological record provides a unique glimpse into the lives (and 45 deaths) of our ancestors. Investigations typically focus on the recording, recovery and analysis of the 46 visible human remains, grave goods and burial structures (Brothwell, 1981). Although some well-47 preserved burials have provided evidence of clothing (Barfield, 1994; Hadian et al., 2012), organic 48 grave goods (Barfield, 1994) and body tissues (O'Connor et al., 2011; Stead et al., 1986), the survival 49 of perishable materials in a condition that can be visually recognised is rare and generally limited to 50 particular burial environments (e.g. waterlogged, arid, frozen) and modes of preparation of the corpse (e.g. embalming and mummification). While organic materials were undoubtedly significant, 51 52 widespread and varied components of human burials, in the vast majority of burials the extent of 53 decay is such that evidence of these components is invisible to traditional archaeological techniques. 54 Organic chemical analysis is now an established tool in the field of archaeology and has been applied

55 to many different types of archaeological remains (Evershed, 2008), including human remains. Bone

56 collagen can survive over archaeological timescales and is routinely isolated for stable isotope and 57 radiocarbon analysis. Exceptional examples of soft tissue preservation have provided rare 58 opportunities to study their protein and lipid compositions and thereby the extent and mechanisms 59 of preservation (Evershed and Connolly, 1988; Gulaçar et al., 1990; Mayer et al., 1997; O'Connor et 60 al., 2011). Likewise, chemical characterisation of organic residues associated with the wrappings of 61 Egyptian mummies has provided information on the preparations used by ancient embalmers (Buckley and Evershed, 2001). By contrast, the potential for the soil in contact with buried human remains to 62 retain molecular information relating to decayed organic components of the burials has received little 63 attention. Such organic remains could relate to components from the body tissues of the interred, 64 from the stomach and gut contents, from formulations used in preparing the body for burial and from 65 66 substances associated with grave goods and burial structures, such as a coffin. Each of these sources 67 has potential to inform the archaeological interpretation: components established to derive from the 68 body tissues, whether or not the physical remains survive, could provide opportunities for isotopic 69 analysis and hence consideration of differences in diet among past populations (Stott et al., 1999); 70 stomach contents can reveal information regarding the balance of plant vs meat in the meals 71 immediately prior to death (Pickering et al., 2014); formulations used in burial rituals could reveal 72 commonalities and differences among individuals within localities and over time; signatures from 73 grave goods could reveal the nature of now decayed materials that were placed within the graves and 74 signatures of the burial structures (Burns et al., 2017) could also reveal commonalities and differences 75 over time and with geographical location. In order for organic signatures to be meaningfully employed 76 it is necessary to gain insights into the nature and extent of their alteration in the environments 77 particular to graves. Whilst much variation can be expected according to burial practice, soil 78 composition and chemistry, hydrology and age of burial, some generalities can be anticipated owing 79 to the presence of a substantial amount of organic matter in a somewhat defined space.

Grave soils from contemporary human burials have been shown to contain chemical signatures of degraded adipose tissue (Bull et al., 2009; Forbes et al., 2002). Given that the recalcitrance of lipids allows their survival over geological timescales (Eigenbrode et al., 2008), informative chemical signatures from wide variety of sources have the potential to be preserved in archaeological grave soils. We present results of the chemical and micromorphological analysis of grave soils and sediments from two human burials from Hungate, York (UK), one of Roman age and the other of Anglo-Scandinavian age.

87 2 Experimental

88 2.1 Excavation

The Roman age grave (C51364, 1st – 4th C AD) and Anglo-Scandinavian age grave (C53700, ¹⁴C date AD 870 - 980) were sampled for the InterArChive project in 2010 and 2011 during excavation of the Hungate site. Both were undisturbed burials, containing articulated skeletons (Fig. 1). The graves presented differences in the levels of preservation, both in terms of completeness of the skeletal remains and in the physical condition of the bones. Whereas the bone in the Anglo-Scandinavian burial was very well preserved, that in the Roman burial had lost much structural integrity and was incomplete, particularly the upper left side of the remains (Fig. 1).





Figure 1. Left panel C53700: Anglo-Scandinavian age burial (AD 870-980) containing the skeletal
 remains of an adult. During excavation, degraded wood was identified believed to have once been a
 wooden lid covering the remains. Right panel C51364: poorly preserved skeletal remains of an adult
 (Roman age ~ 3rd century AD) and a nearby Anglo-Scandinavian age cess pit (Context 2652) that cuts
 the grave.

101 2.2 Sampling

102 Controls were collected from non-grave soil from the site (C1) and from grave fill above the level of 103 the skeletal remains, C2 and C3, as essential comparators of grave fill that has not been in contact 104 with the buried remains (Fig. 2). 105 Samples of the burial matrix were collected in line with the InterArChive sampling strategy (Usai et al., 106 2014) from key anatomical locations on the skeleton (Table 1). The maximum number of prescribed 107 points around the skeletal remains were targeted for chemical analysis and 4 key areas; head (1), 108 pelvis (2) and feet (3/4) were targeted for micromorphological analysis (Table 1). Additional samples 109 (prefixed by the letter 'A') were collected in response to specific features of the individual graves and 110 from a cess pit adjacent to the Roman burial. Samples for micromorphology were collected from undisturbed archaeological sediment or grave fills using Kubiena tins (83 × 50 × 32 mm, 83 × 55 × 42 111 112 mm, and $65 \times 38 \times 28$ mm). On return to the laboratory, samples were refrigerated prior to thin 113 section preparation. Samples for chemical analysis were collected proximal to the skeletal remains or feature using a spatula, wrapped in pre-cleaned foil, placed in geochemical sample bags and stored 114 115 cold. On return to the laboratory, samples were stored frozen until required for analysis.

116 2.3 Manufacture of thin sections

The samples were dried in the tins through acetone vapour exchange (Miedema et al., 1974) and impregnated under vacuum with a slow curing polyester resin (Polylite 32320-00). The resulting consolidated soil blocks were cut to produce the thin sections. The cut sections were back-polished, mounted, cut and ground to 30 µm thickness, with final 3 µm and 1 µm polishes.

121 2.4 Optical microscopy

Micromorphological observations were performed using a petrographic microscope (Zeiss AxioLab A1) equipped with a Zeiss AxioCamERc5s camera and AxioVision imaging software. The standard terminologies proposed by Bullock et al. and by Stoops (Bullock et al., 1985; Stoops, 2003) were adopted for the descriptions of the slides and semi-quantitative estimation of the features.

126 2.5 Scanning electron microscopy-energy dispersive X-ray analysis (SEM-EDS)

127 SEM-EDS analyses were conducted at the University of Stirling employing an EVO MA 15 Zeiss scanning 128 electron microscope with workflow automation and an Oxford Instruments INCA X-Max EDS to 129 provide micro-chemical data of fine material and inorganic pedofeatures of uncertain interpretation 130 (Courty et al., 1989; Fitzpatrick, 1993; Goldberg and Macphail, 2003). Standard operating conditions were 8.5 mm working distance and 20 kV accelerating voltage. Calibration was achieved through the 131 132 analysis of standard cobalt every 2 h and standard dolomite at the beginning of each session. A 133 minimum of seven individual point analyses were carried out for each measured region. The 134 measurements of O and C were excluded owing to the presence of the resin in which the samples 135 were consolidated and weight percent data were normalized to 100%.

136

Anglo-Scandinavian burial C53700		Roman burial C51364				
Sample Number [†]		Position	Sample Number ⁺		Position	
C1		Control (non-grave soil)	C2	М	Control (upper grave fill)	
C2		Control (lower grave fill)	С3	М	Control (lower grave fill leve	
					with top of skull)	
C3		Control (lower grave fill)	1	М	Skull	
1	М	Skull	2	М	Sacrum	
1b		Below skull	3		Left foot	
2	М	Sacrum	4		Right foot	
2b		Below sacrum	3 & 4	М	Between feet	
3		Both feet	5		Left shoulder	
5		Below left shoulder	6		Right shoulder	
6		Below right shoulder	7		Left elbow	
7		Below left elbow	8		Right elbow	
8		Below right elbow	9		Left hip	
9		Below left hip	10		Right hip	
10		Below right hip	11		Left knee	
11		Below left knee	12		Right knee	
12		Below right knee	13		Left iliac	
13		Below left iliac	14		Right iliac	
14		Below right iliac	15		Sternum	
15		Below sternum	16		Right hand	
A1		Right scapula	A1		Wall of cess pit (context 2652)	
A2		Dark soil right pelvis	A2		Material excavated from cess pit	
					(context 2652)	
A3		Соссух	A3		Intracranial sediment	
			A4		Abdominal region	
			A5		Dark soil within pelvis	

137 **Table 1**. Soil samples collected for the two graves from Hungate, York, UK.

¹Sample number corresponds to the standard sampling locations defined in Usai et al., (2014). Unless specified otherwise the soil adjacent
 to the bone was collected. Samples prefixed by the letter A represent samples additional to the standard sampling locations. M indicates

100 to the bone was concluded. Samples prenzed by the retter A represent samples additional to the standard sampling locations. In inc

140 that an undisturbed block sample for micromorphological analysis was collected in addition to a loose soil sample.

141 2.6 Preparation of materials and extracts for chemical analysis

All solvents used were analytical grade or higher. All glassware was baked (450°C, 6 h) prior to use 142 143 using a Pyroclean Trace oven (Barnstead International, USA) to remove organic contaminants. Frozen 144 soil samples were freeze dried using a HETO PowerDry PL3000 (Thermo, Hemel Hempstead, UK). Dry 145 soils were disaggregated with a pestle and mortar, and sieved using 1 mm, 400 μ m and 200 μ m sieves. 146 All subsequent work was performed on the sub 200 µm fraction. Extraction was performed using an 147 accelerated solvent extraction system (ASE 350, Dionex, Hemel Hempstead, UK). Prior to loading soil 148 samples, the empty extraction cells were subjected to extraction using the same solvents and conditions as for the samples. Soil (3-6 g) was loaded into 5 ml stainless steel cells and extracted three 149

150 times with dichloromethane: methanol (9:1 v/v; 5 min at 100°C and 1500 psi). A blank extraction was 151 also performed to assess whether any contamination was being introduced at the extraction step. 152 Solvent was removed using a rotary vacuum concentrator (RVC 2-25, Christ, Osterode am Harz, DE). 153 The total extract was divided into two portions, one of which was fractionated using a custom-made 154 glass chromatography column (90 mm x 10 mm i.d.) packed with silica gel (750 mg, height in column 155 = 20 mm). The column was conditioned by washing with three bed volumes each of dichloromethane:methanol (1:1, v/v) and dichloromethane followed by equilibration with hexane. 156 157 The extract to be fractionated was dissolved in dichoromethane (200μ l), spiked onto silica gel (40 mg) 158 and the solvent removed in vacuo. The impregnated silica was loaded onto the top of the packed 159 chromatography column. The column was washed with three bed volumes each of i) hexane, ii) 160 hexane:toluene (1:1, v/v), iii) hexane:ethyl acetate (4:1, v/v) and iv) dichloromethane:methanol (1:1, 161 v/v) to generate four chromatographic fractions: i) aliphatic hydrocarbons, ii) aromatic hydrocarbons 162 and *n*-alkanals, iii) alcohols and esters and iv) acids, respectively. Solvent was removed using a rotary 163 vacuum concentrator.

164 2.7 Derivatisation

Total extracts and polar (iii and iv) fractions were reconstituted in dichloromethane:methanol (2:1, v/v; 300 μ l) followed by addition of trimethylsilyl diazomethane (20 μ l) and allowed to react for 30 min, before drying under a gentle stream of nitrogen gas. Immediately prior to analysis, total extracts and polar fractions were heated with *N*,*O*-bis(trimethylsilyl)trifluoroacetamide (100 μ l, containing 1% trimethylchlorosilane) and 5 drops of pyridine for 1.5 h at 60°C before removing excess derivatising agent under a gentle stream of nitrogen gas.

171 2.8 Total organic carbon (TOC) analysis

Total organic carbon (TOC) analysis was performed using a Flash 2000 elemental analyser (Thermo, 172 173 Hemel Hempstead, UK) fitted with a MAS200R autosampler, chromatographic column and thermal 174 conductivity detector. Helium was used as a carrier gas at a flow rate of 140 ml min⁻¹. Soil samples 175 (10-15 mg) were weighed using a XS3DU microbalance (Mettler Toledo, Leicester, UK) into silver foil 176 capsules. Samples were treated with 2 drops of hydrochloric acid (6 M) to destroy carbonates before 177 heating to 80°C for 6 min to drive off excess acid solution. Foil capsules were folded to exclude air. 178 Samples were combusted in a quartz reactor tube packed with copper oxide granules and electrolytic 179 copper wires held at 900°C. Sample introduction coincided with a pulse of oxygen (250 ml min⁻¹, 5 s). 180 Instrument control, data acquisition and processing was by Eager Xperience V1.11 (Thermo, Hemel 181 Hempstead, UK).

182 2.9 Gas chromatography (GC) and gas chromatography-mass spectrometry (GC-MS)

183 Total lipid extracts were analysed using a Trace GC Ultra gas chromatograph (Thermo, Hemel 184 Hempstead, UK) equipped with a Triplus autosampler, a fused silica capillary column (J&W DB-5, 60 185 m x 0.32 mm i.d., 0.25 µm film thickness, Agilent, Wokingham, UK) and a flame ionisation detector 186 (FID). Samples were prepared in dichloromethane and volumes of 1 μ l injected onto the column via 187 a split/splitless injection port operating in splitless mode (280°C, split flow 1:25, splitless time 0.8 min). 188 The oven temperature was programmed from an initial temperature of 70°C to 130°C at a rate of 20°C 189 min⁻¹ and then to 320°C at a rate of 4°C min⁻¹ where it was held for 40 min. Helium was used as a 190 carrier gas at a flow rate of 2 ml min⁻¹. The FID was held at 330°C. Instrument control, data acquisition 191 and analysis was by ChromQuest 5.0 V3.2.1 (Thermo, Hemel Hempstead, UK).

192 Chromatographic fractions (i-iv) were analysed using an identical gas chromatograph to that above 193 equipped with an ultrafast column module (UFC-5, 10 m x 0.1 mm i.d., 0.1 μ m film thickness, Thermo, 194 Hemel Hempstead, UK). Samples were prepared in dichloromethane for analysis and volumes of 1 μ l 195 injected onto the column via a split/splitless injection port operating in split mode (280°C, split flow 1100) and the FID held at 330°C. The column oven was programmed from an initial temperature of 197 50°C (0.5 min isothermal) to 330°C at a rate of 90°C min⁻¹ where it was held for 4 min. Helium was 198 used as the carrier gas at a flow rate of 0.5 ml min⁻¹.

199 GC-MS analysis was performed on selected fractions using a 7860A gas chromatograph (Agilent, 200 Wokingham, UK) equipped with a 7683B Series autosampler coupled to a GCT Premier Micromass 201 time of flight mass spectrometer (Waters, Elstree, UK). Separation was achieved on a fused silica 202 capillary column (Zebron, ZB-5, 30 m x 0.25 mm i.d., 0.25 µm film thickness, Phenomenex, 203 Macclesfield, UK). Samples were prepared in dichloromethane for analysis and volumes of 1 μ l 204 injected onto the column via a split/splitless injection port operating in split mode (280°C, split flow 205 1:5). The oven was programmed from an initial temperature of 70°C to 130°C at a rate of 20°C min⁻¹ and then to 320°C at a rate of 4°C min⁻¹ where it was held for 40 min. Helium was used as the carrier 206 207 gas at a flow of 1 ml min⁻¹ and the MS transfer line set to 300°C. Mass spectra were acquired over the 208 range m/z 50-750 (cycle time 0.2 s) using an electron ionisation energy of 70 eV. Instrument control, 209 data acquisition and processing was by MassLynx V4.1 (Waters, Elstree, UK). Compounds were 210 identified from interpretation of their mass spectra and comparison with library spectra (NIST 08) 211 where available.

212 2.10 High performance liquid chromatography-multistage tandem mass spectrometry (HPLC-MSⁿ)

HPLC-MSⁿ analysis of triacylglycerols was performed on selected total extracts according to Method J
 of (Hasan, 2010) using a Dionex Ultimate 3000 rapid separation liquid chromatograph, controlled by

215 Chromeleon 6.8 (Dionex, Hemel Hempstead, UK), coupled to a HCTultra ETD II quadrupole ion trap 216 mass spectrometer (Bruker Daltonics; Coventry, UK) fitted with an atmospheric chemical ionisation 217 (APCI) source. Briefly, samples were prepared in hexane:propan-2-ol:acetonitrile (1:1:1, v/v/v). Lipid 218 separation was accomplished using two 3 µm Spherisorb ODS2 columns (150 x 4.6 mm i.d.) coupled 219 in series, using a ternary solvent gradient system comprising acetonitrile, dichloromethane and 220 ammonium acetate in methanol (0.01 M) at a flow rate of 1 ml min⁻¹. APCI was operated in positive 221 ion mode using the following settings: vaporiser temperature 450°C, nebuliser gas (N₂) 50 psi, drying 222 gas (N₂) flow rate 3 L min⁻¹, drying gas temperature 150°C, corona 4000 nA and capillary voltage -4000 223 V. Mass spectra were acquired in Ultra Scan mode over the range m/z 300-950 with a cycle time of 224 0.025 s. Online multistage tandem mass spectra (to MS³) were generated following selection of the 225 most abundant ion in the mass spectrum for collision induced dissociation. The isolation width was 226 set to 4 *m/z* units, maximum accumulation time was fixed at 200 ms and Smart Fragmentation set to 227 between 30-200% of a fragmentation amplitude of 1.0 V. MS control, data acquisition and processing 228 was by Compass EsquireControl 6.2 and Compass DataAnalysis 4.0 (Bruker Daltonics; Coventry, UK).

229 3 Results and Discussion

230 Chemical and micromorphological analyses of samples of burial soil, collected from numerous specific 231 locations (e.g. Fig. 2; full sample list in Table 1) around the skeletal remains of the Roman and Anglo-232 Scandinavian burials, were compared with controls to identify features and signatures uniquely 233 associated with the burials. The soil particles were dominated by coarse grains of quartz with clays dominating the fine material (Table 2). The soil in the Roman burial was apedal whereas that in the 234 235 Anglo-Scandinavian burial contained weakly developed granular peds and more voids. Given their 236 different positions within the stratigraphic profile it is not possible to determine if the difference 237 between the two graves relates to intrinsic differences during soil formation or to post depositional 238 disaggregation of peds in the Roman grave. The sediment from both graves contained fragments of 239 humified plant matter including gymnosperm wood in the Anglo-Scandinavian grave. Fungal hyphae 240 were also identified in the Anglo-Scandinavian grave in the area of the pelvis and sclerotia in the areas 241 of the skull and pelvis. The dominant pedofeatures in the Roman grave were spherulites, goethite 242 crystals, Fe/Mn nodules and clay coatings indicating variations in redox conditions and water 243 saturation. Amorphous phosphate was absent or very low around the torso in the Roman grave and 244 the concentrations at the feet and in the controls were anomalously low compared with most other graves examined in Hungate (Ghislandi, 2016). The Anglo-Scandinavian grave contained higher levels 245 246 of amorphous phosphate throughout the thin sections and vivianite in association with the bone. 247 Hence, parts of the Roman grave may have been subjected to processes that caused leaching of phosphorus. The presence of Fe/Mn nodules in the Anglo-Scandinavian grave indicate changes in
water saturation, though in a generally more oxidising environment than in the Roman grave.

250 The organic extracts of the soils were analysed by gas chromatography-mass spectrometry (GC-MS) 251 and high performance liquid chromatography-multistage tandem mass spectrometry (HPLC-MSⁿ) to 252 determine the distributions of lipid components in the soils. The extracts of all samples and controls from both graves were dominated by autochthonous components of soil organic matter, mainly 253 254 homologous series of *n*-alkanoic acids (C_{16} - C_{34}), *n*-alkanols (C_{18} - C_{34}) and *n*-alkanes (C_{20} - C_{35}), the 255 distributions being characteristic of the leaf waxes of vascular plants (Eglinton and Hamilton, 1967). 256 In addition, phytosterols 24-ethylcholest-5-en-3β-ol (sitosterol), 24-ethylcholest-5,22-dien-3β-ol 257 (stigmasterol) and 24-methylcholest-5-en-3 β -ol (campesterol) were detected.

258 Despite the strong overprint from the natural soil organic matter, the samples associated with the 259 skeletal remains exhibited distinct differences to the controls, generally exhibiting elevated levels of 260 triacylglycerols (TAGs) (e.g. Figs. 2a and 3); di- $C_{16:0}$ and $C_{16:0}/C_{18:0}$ 1,2- and 1,3-diacylglycerols (DAGs); C16:0 and C18:0 monoacylglycerols (MAGs); C16:0 and C18:0 fatty acids (Fig. 2b); C16:0 and C18:0 n-alkan-1-als 261 262 (Fig. 2b) and branched chain iso (i-) and anteiso (ai-) C_{15:0} and C_{17:0} fatty acids (Fig. 2c). The TAG 263 concentrations associated with the Anglo-Scandinavian remains were around an order of magnitude 264 greater than those for the Roman remains. Though TAGs have been identified from within the 265 protective environment of archaeological ceramics (Garnier et al., 2009; Kimpe et al., 2001; Mirabaud 266 et al., 2007; Šoberl et al., 2008) their occurrence in grave soils has not been reported previously. The 267 elevated levels of the TAGs, the primary constituents of animal fats and vegetable oils, associated with 268 the skeletal remains suggests they derive from adipose tissue. The low levels associated with the 269 controls (e.q. Fig. 2) may reflect minor contributions from plant sources or background levels from 270 previous burials within the cemetery. Notably, the TAG distributions in the soil samples are dominated 271 by saturated components (Fig. 3b) whereas fresh human adipose tissue comprises predominantly 272 (99%) unsaturated TAGs (Mayer et al., 1997). Generally, unsaturated components are more 273 susceptible to degradation (e.g. via oxidative cleavage) than their saturated counterparts. In addition, 274 human lipases have been shown to exhibit a preference for acyl moieties at the *sn*-1 and *sn*-3 positions 275 on the glycerol backbone as well as for unsaturated moieties (Jensen et al., 1994; Raclot and Groscolas, 276 1993). Preferential degradation of TAGs containing unsaturated acyl chains could explain the 277 observed predominance of saturated TAGs in the soil extracts. Comparison of the acyl chain 278 distributions of the saturated TAGs in the soil extracts from around the skeletal remains (Fig. 3c) with 279 the saturated TAG-derived fatty acid distributions of fresh human adipose tissue (Hodson et al., 2008) 280 reveals a strong similarity (A; Fig. 3c). By contrast, complete reduction of monounsaturated and of all 281 unsaturated fatty acids of human adipose TAGs would generate the predicted ratios shown for B and

282 C (Fig. 3c), respectively. Thus, extensive post-depositional alteration of the body fats is evident and 283 involves preferential loss of unsaturated TAG moieties. The TAG distributions in the soils around the 284 skeletal remains (Fig. 3b) exhibit a striking similarity to those reported for the adipocere of a 65 year 285 old ice mummy and the skin of a 5300 year old ice mummy (Mayer et al., 1997), supporting the genesis 286 of these signatures from human adipose tissue. In those cases, preservation of TAGs may be 287 attributed largely to the physiochemical barriers to degradation: encasement in ice and desiccation, respectively. A study involving augmentation of three different soils with a model saturated TAG 288 289 (tristearin) showed that release of fatty acids within the first weeks of incubation at 20°C could be 290 attributed to the soil microbial community (Hita et al., 1996). It is remarkable, therefore, that TAG 291 signatures of adipose tissue can be recovered from grave soils after more than 1600 years since 292 interment. TAGs and cell membrane phospholipids are degraded in soils by stepwise hydrolysis to 293 fatty acids and glycerol (and phosphate in the case of phospholipids), which explains the elevated 294 levels of DAGs, MAGs and C_{16:0} and C_{18:0} fatty acids compared with the controls (e.g. Fig. 2b).

295 The $C_{16:0}$ and $C_{18:0}$ *n*-alkan-1-als associated with the skeletal remains have not been reported as 296 significant components of soils and were not detected in the controls (Fig. 2b). Short chain n-alkan-1-297 als, principally C₁₆ and C₁₈ components, occur as both free and plasmalogen-bound components in various animal tissues (Gilbertson et al., 1967; Rapport and Lerner, 1959; Wittenberg et al., 1956) 298 299 although their amounts are too low (relative to other lipid classes such as fatty acids) to represent a 300 significant source of *n*-alkan-1-als in the soil extracts, where they represent approximately 10% 301 abundance relative to summed $C_{16:0}$ and $C_{18:0}$ fatty acids (Fig. 2b). The most likely source that would 302 explain the concentrations of *n*-alkan-1-als in the soil extracts is reduction of fatty acids (e.g. mediated 303 by anaerobic bacteria). The concentrations of *n*-alkan-1-als in the soils are not, however, controlled 304 by fatty acid concentration alone: the differences in C_{16:0} and C_{18:0} fatty acids concentrations with 305 anatomical location do not correlate with those of the corresponding n-alkan-1-als (Fig. 2b), most 306 notably around the pelvis in the Anglo-Scandinavian grave. Features of bone morphology (for example 307 the skull or sacrum) could restrict water movement and contribute to highly localised regions of anoxia 308 that favour reduction of organic matter. Thus, the reductive transformation of fatty acids to *n*-alkan-309 1-als, and the survival of the relatively labile aldehyde species themselves, is strongly indicative of 310 persistent oxygen-limited conditions within areas of the grave, with higher n-alkanal:fatty acid ratios 311 reflecting localised regions of more extensive anoxia.

Table 2a. Micromorphological descriptions of the sediments from the Roman age grave (C51364) and Scandinavian age grave (C53700) from Hungate, York,

313 UK). PPL = plane polarised light, XPL = cross polarised light.

Burial	Microstructure	Voids	Mineral	Organic components	Pedofeatures
			components		
C51364	Two principal c/f related	Chambers, 500-2000 μm	Mineral	Fragments of weathered	Sideritic spherulites in the areas of skull, pelvis, feet and C3 exhibit
(Roman)	distributions either strongly	with undulating	components mostly	or partially weathered	morphological differences: Group 1 (skull 5-10%, 5-12 μm; feet 10-20%,
	separated or intermingled:	surfaces, in the areas of	sub-angular and not	humified plant	2-5 μm ; pelvis 2-5%, 5-12 μm): orange in PPL; Maltese cross extinction in
	porphyric and	the feet (8%), skull and	weathered quartz	structures in all samples	XPL. Infillings in pores and within fine material in the areas of skull and
	poorly/moderately sorted	C3 (5%) and pelvis (2%).	(45-55%).	(2%) except control C3.	pelvis. Sealed between well-preserved layered coatings of Fe rich fine
	(10-30% by area) or chitonic	Channels at all positions	Occasional	Plant remains,	material in area of feet. Group 2 (C3, skull and pelvis): 5-20 μm ; yellow in
	and moderately/well sorted	(5% except for skull:	fragments of	50-2000 μm,	PPL, amorphous in XPL; infillings in pores and among mineral grains. Some
	(70-90% by area). Several	8%).	guartzite in the	brown/black in PPL and	have coalesced to form dumbbell structures. Group 3 (feet): 5-15 $\mu\text{m};$
	intermediate distributions.	0,01.	areas of the feet	isotropic in XPL. Some	orange/yellow in PPL. Some present two distinct layers with high contrast
	Porphyric areas were	Modified complex voids	(2%) and C2 (5%).	coarser fragments	at junction; only outer layer exhibits Maltese cross extinction; core
	brown/light brown/yellowish	only in the areas of skull	(270) and C2 (370).	exhibit vessels, hence	amorphous in XPL.
	brown in PPL and	and pelvis (5%) and	Rare grains of	angiosperm-derived.	
	brown/yellow or isotropic in	control C2 (8%).	biotite were in the		Infillings of goethite in voids in area of skull (8%): loose discontinuous or
	XPL, with dotted limpidity and	Frequency of packing	area of the skull		dense incomplete. Occurs as fan-like crystals, orange/red colour in PPL and
	speckled b-fabric. Chitonic	voids increased from the	(2%).		reddish brown or isotropic in XPL and with parallel striated b-fabric.
	areas were yellowish or	skull to the control C2,			Sub-angular/sub-rounded Fe/Mn nodules (50-500 $\mu m,$ 5%) in all samples
	reddish brown in PPL and	absent from the pelvis.			except control C2.
	orange/yellow or isotropic in	Planar voids, thickness			Thin dark brown coating around voids, probably organic in origin, in areas
	XPL, with speckled or parallel	100-1500 μm, in areas of			of feet (8%), skull and pelvis (5%). External quasicoatings of clay around
	striated b-fabric.	skull (15%), pelvis and			voids in area of pelvis (8%) and C2 (5%), clay textural impregnations in area
	Samples were apedal.	feet (5%) and C3 (8%).			of skull. Coatings of amorphous phosphate in control samples (5-10%).

Table 2b. Micromorphological descriptions of the sediments from the Roman age grave (C51364) and Scandinavian age grave (C53700) from Hungate, York, UK).

Burial	Microstructure	Voids	Mineral components	Organic components	Pedofeatures
C53700	Two different fabrics: A (50-85%), B (20-50%)	Chambers in areas of	Quartz dominant in all	Amorphous organic matter in area of skull	Amorphous phosphates in all samples,
(Anglo-	and C (10%) in varying proportions. Type A:	pelvis (8-15%) and skull	samples (30-50%),	(2%). Humified plant structures (200-	most frequent in area of pelvis (5-20%).
Scandinavian)	close porphyric c/f related distribution, good	(2%).	quartzite only in areas	2000 $\mu m)$ in all samples (2-5%).	Occur as coatings voids, nodules and
	sorting. Fine material, 10-40%, yellowish	Channels in areas of	of pelvis (2%) and	Occasional fragments characterized by	strong impregnations in fine material;
	brown/brown in PPL and reddish brown or	pelvis (8-15%) and	vertebrae (5%). Flint	vessels, hence angiosperm-derived.	light yellow or yellow in PPL and
	yellow/brown in XPL, with dotted limpidity	vertebrae (10%).	in area of pelvis (2%).	Gymnosperm wood indicated by single	isotropic in XPL; some exhibit crystalliti
	and speckled b-fabric. Type B: open porphyric	vertebrae (10%).		imprint of fragment in fine material in	layers, occasionally weathered with
	and well sorted. Fine material, 60-80%,	Modified complex voids		area of pelvis. Same sample exhibits	granular aspect in pelvic area.
	reddish brown or yellowish brown in PPL and	in areas of skull (8-25%),		humified plant (>2000 μ m) comprising	Occasional amorphous phosphate
	yellow/brown in XPL, dotted limpidity and	vertebrae and pelvis		four fragments oriented parallel to the	coatings around spherical Mn nodules
	speckled or parallel striated b-fabric. Type C	(8%).		resting plane, circular/angular brown cells	areas of skull and pelvis. Single
	(area of vertebrae): chitonic c/f related	Packing voids in the		with infilling of fine material. Light brown	occurrence of vivianite (100-150 μ m) i
	distribution, good sorting, yellow/orange in	areas of the skull (8%),		in PPL and brown or isotropic in XPL.	area of pelvis.
	PPL, yellow or isotropic in XPL, limpid or	pelvis (5%) and		Fungal hyphae present in groundmass in	Red redox impregnations in areas of
	dotted limpidity, parallel striated b-fabric.	vertebrae (15%).		area of pelvis. Sclerotia (20-200 μ m) in	pelvis and vertebrae (5%). Fe/Mn
	Samples from areas of skull and pelvis exhibit			areas of skull and pelvis (2%).	nodules in area of skull (2-8%) and
	spongy appearing microstructures.				
	Weakly daysland granular pads (E. 10%) in			Occasional roots in area of pelvis: orange	occasional in the other samples.
	Weakly developed granular peds (5-10%) in			in PPL and partially white in the internal	Clay coatings around mineral grains in
	the area of vertebrae: fine in size and			section in XPL.	areas of skull (2%) and pelvis (5%).
	unaccommodated.				

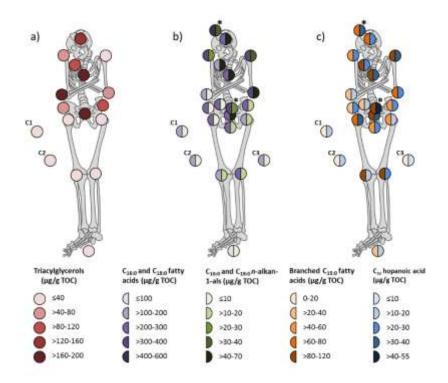
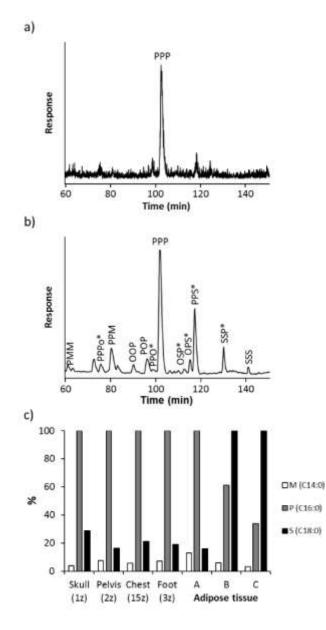




Figure 2. Variation in concentrations of specific lipids in the Anglo-Scandinavian age burial C53700: a) combined triacylglycerols (TAGs; red shades), b) combined $C_{16:0}$ and $C_{18:0}$ fatty acids (purple shades) combined $C_{16:0}$ and $C_{18:0}$ *n*-alkan-1-als (green shades) and c) combined *i*- and *ai*- $C_{15:0}$ fatty acids (orange shades) and bishomohopanoic acid (C_{32} ; blue shades). All soil samples were collected from below the skeletal remains except for samples marked with an asterisk (*) which were collected from adjacent to the skeletal remains. C1 = control from outside the grave cut, C2 = control from fill at the edge of the grave above the level of the skeleton, C3 = control of the grave fill, level with the skeleton.

Branched chain fatty acids (*i*-C_{15:0}, *ai*-C_{15:0}, *i*-C_{17:0}, and *ai*-C_{17:0}) are significant components of bacterial 326 327 lipids while being scarce in other sources. The spatial variation of these components in the Anglo-328 Scandinavian grave (Fig. 2c) was also mirrored by a series of C_{30} - C_{33} hopanoic acids and hopanols (e.g. 329 Fig. 2c), diagenetic products of the bacteriohopanetetrol and/or aminobacteriohopanetriol 330 components of bacterial cell membranes (Quirk et al., 1984; Winkler et al., 2001). The variability in 331 the bacterial markers across the remains (Fig. 2c) may be explained by the development and spread 332 of populations evolved from the gut microbial fauna during decomposition (Can et al., 2014) together 333 with soil microbe invasion in response to the corpse providing a plentiful supply of organic substrate. 334 Anatomical locations where the highest proportions of body tissue are distributed appear to contain 335 more abundant bacterial markers, indicating more intense/protracted microbial activity.



336

Figure 3. Partial HPLC-MS chromatograms (*m*/z 300-950) of the total lipid extracts for a) the control 337 (C1) and b) pelvis (2) from C53700 acquired using Method J (Hasan, 2010). Triacylglycerols (TAGs) are 338 339 labelled with letters to denote the acyl moieties attached to the glycerol backbone e.g. tripalmitin = PPP. M = myristyl ($C_{14:0}$), P = palmityl ($C_{16:0}$), Po = palmitoleyl ($C_{16:1}$), S = stearyl ($C_{18:0}$) and O = oleyl 340 341 (C18:1). *Indicates that only one of several possible positional isomers is shown. c) Comparison of the relative abundances of saturated TAG fatty acids (M = $C_{14:0}$, P = $C_{16:0}$ and S = $C_{18:0}$), normalised to the 342 major component, in the soil samples from around the skeletal remains of C53700 to those of fresh 343 344 human adipose tissue (Hodson et al., 2008). A = relative abundance of saturated fatty acids for human adipose tissue; B = relative abundance of saturated fatty acids for human adipose tissue assuming 345 346 reduction of all monounsaturated fatty acids to saturated components; C = relative abundance of 347 saturated fatty acids for human adipose tissue assuming reduction of all unsaturated fatty acids to 348 saturated components.

349 In both graves the animal sterol cholest-5-en- 3β -ol (cholesterol) was generally present in greater 350 concentration in the samples associated with skeletal remains than in the controls (e.g. Fig. 4a) with 351 the Anglo-Scandinavian burial having the higher concentrations. The preservation state of the Roman 352 skeletal remains, C51364, was extremely poor compared with the Anglo-Scandinavian remains, 353 consistent with its greater age and wetter burial environment. Interestingly, cholesterol 354 concentrations were more varied, being particularly pronounced for soils associated with the more 355 degenerated regions of the skeleton, closest to an Anglo-Scandinavian age cess pit that was proximal 356 to the upper left hand side of the remains at a level slightly above the burial (Fig. 4a). Identification 357 of cholesterol as a significant lipid component of archaeological bone (Evershed et al., 1995; Jim et al., 358 2004) together with the degenerated skeletal remains indicate it to have been released into the soil 359 as the bone degraded. Hence, persistent cholesterol signatures may have value in aiding the 360 identification of the original location of human remains in graves where the bones have not survived, 361 subject to appropriate caution being exercised to differentiate signatures from alternative sources 362 such as soil fauna. The cholesterol in the extracts was accompanied by its reduction products 5α -363 cholestan-3 β -ol and 5 β -cholestan-3 β -ol (coprostanol). 5 α -Cholestan-3 β -ol is the major product of 364 microbial reduction of cholesterol in soils and plant and mammalian tissues (Bethell et al., 1994). Its 365 epimer, 5 β -cholestan-3 β -ol, typically a minor product of cholesterol reduction in the environment, is 366 formed in significant amounts during microbial transformation of cholesterol in the guts of most 367 higher animals and is the major sterol in human faecal material (Leeming et al., 1996). Accordingly, the ratio of $5\beta/(5\alpha + 5\beta)$ cholestanols has been used as a proxy to indicate faecal contamination of 368 369 sediments and water systems: values >0.7 indicate pollution (Grimalt et al., 1990). For the Anglo-370 Scandinavian grave, $5\beta/(5\alpha + 5\beta)$ cholestanol index values suggestive of faecal contamination were 371 restricted to the samples from the pelvic region, signifying inputs from the gastrointestinal tract. For 372 the Roman grave, however, high ratios were observed throughout the burial (Fig. 4b), revealing clear 373 evidence for the ingress from the cess pit of material having a strong faecal signature. The high $5\beta/(5\alpha)$ 374 + 5 β) cholestanol index values of samples from the cess pit combined with their lower cholesterol 375 levels than the closest samples from within the grave fully support the interpretations discussed 376 above. Despite the influence of the cess pit, the organic signature from the pelvic region of C51364 377 can still be distinguished by abundant bacterially-derived C₃₀-C₃₃ hopanoic acids (Fig. 4c). 378 Furthermore, hopanoic acid distributions at the feet suggest migration of gut-derived organic matter 379 within the coffin. The evidence for organic signatures derived from the cess pit leachate affecting the 380 preservation of the bones comes from the generally worse preservation of the skeletal remains nearer 381 to the pit. Hence ingress of leachate from the Anglo-Scandinavian age cess pit into the grave 382 contributed a strong faecal signature to the soils of the upper left section of the remains. Such organic

rich leachate would have been strongly anoxic and acidic, the latter accounting for the decalcification 383 384 of the bone and consequent liberation of cholesterol within its area of influence. The acidic and 385 reducing nature of the leachate means that it would have also carried ferrous iron in solution and is 386 likely to have played a significant role in controlling phosphate levels in the Roman grave, mobilisation 387 of phosphate being most prevalent in mildly acidic solutions (Nriagu, 1972). The absence/low levels 388 of phosphorus in the burial soils of the Roman grave may be due either to low input to the original 389 sediment or, given the greater age of the Roman burial, to the natural decrease in phosphorus levels 390 with soil weathering (Cross and Schlesinger, 1995).

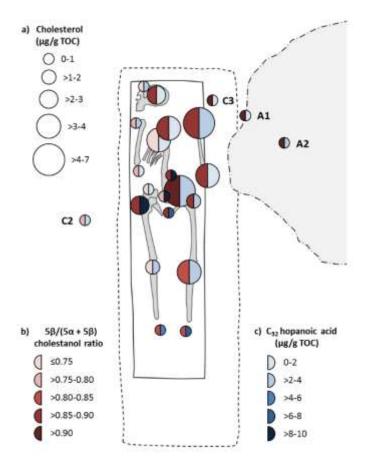


Figure 4. Variation in concentrations of specific lipids in Roman age burial C51364: a) cholesterol (circle sizes), b) 5β-cholestanol/(5α-cholestanol + 5β-cholestanol) ratio (red shades) and c) bishomohopanoic acid (C₃₂; blue shades). C2 = control from natural soil of grave cut above skeleton, C3 = control of grave fill level with top of skull, A1 = soil from natural soil of cess pit cut, A2 = soil excavated from the cess pit. The solid line outlines the soil stain from a coffin, the dashed line outlines the grave cut and the dashed and dotted line marks the boundary of an Anglo-Scandinavian age cess pit that cuts the grave.

398 Micromorphological analysis of undisturbed soil from control C3 and the head, pelvis and feet of the 399 Roman burial revealed spherulites, features absent from control C2 and from all samples from the 400 Anglo-Scandinavian burial. Morphological differences allowed the spherulites to be classified into 401 three distinct groups (Fig. 5, Table 2a). Group 1 spherulites: were mostly present near the skull and 402 feet, and were less abundant near the pelvis. They occurred as infillings in the pores and within the 403 fine material in the areas of the skull and pelvis with intermediate size range (Figs 5 and 6). Around 404 the feet a smaller and narrower size range was present sealed between well preserved layered 405 coatings of Fe rich fine material. Group 2 spherulites were present as infillings in pores and among 406 mineral grains. They were most abundant in C3 and around the pelvis, with lower levels present in 407 the skull area and only traces at the feet (Fig. 5). In some cases, the spherulites had coalesced to form 408 dumbbell structures (Fig, 6a). Group 3 was identified only around the pelvis; some presented two 409 distinct layers with high contrast at their junction with only the outer part of the spherule showing 410 Maltese cross extinction and the core appearing amorphous in XPL.

Designation	Morphology and optical characteristics PPL XPL	Size /μm	Frequency /%	Locatio n
Group 1	-	5-12	5-10	skull
			2-5	pelvis
		2-10	10-20	feet
Group 2		5-20	10-20	C3
		1	2-5	skull
			10-20	pelvis
			<2-5	feet
Group 3		5-15	5-10	pelvis

411 **Figure 5**. Key characteristics and occurrences of different forms of siderite present in the Roman

412 grave from Hungate, York. Detailed of the different spherulite characteristics are given in Table 2a.

413 Scanning electron microscopy-energy dispersive X-ray spectroscopy (SEM-EDS) analysis revealed the 414 spherulites to have appreciable Fe contents. Taking account of literature reports (Driese et al., 2010; 415 Ludvigson et al., 2013; Pye et al., 1990), the composition is attributed to siderite (FeCO₃). Pedogenic 416 siderite is characteristic of wet and persistent saturated soils (Browne and Kingston, 1993), often in 417 association with vivianite and more rarely with goethite, where it occurs as small crystallites or 418 spherulitic aggregates in the groundmass or as pore infillings (Stoops and Delvigne, 1990). It forms as 419 a colourless mineral in reducing conditions, remaining stable at neutral to basic pH and can only 420 precipitate from weakly acid solution if the concentration of dissolved iron is abnormally high (Lemos

421 et al., 2007). Siderite spherules have been observed in organic-rich archaeological floor deposits, in 422 areas inhabited by livestock where reducing conditions were inferred (Gebhardt and Langohr, 1999; 423 Milek, 2012). The Fe contents of the Group 1 spherulites were greater than those of Group 2, with 424 those of Group 1 from the area of the feet also having higher levels of Ca. Group 3 spherulites were 425 characterized by an Fe rich corona with significantly lower Fe levels in the core. This difference is 426 evident in the SEM-EDS images as a bright outer ring and a dull centre (Fig. 6d). The spherulites of the 427 first group showed the highest reflectivity. Sideritic spherulites were observed in several other graves 428 from Hungate, though exclusively in the sediments surrounding the skeletons and not in the upper 429 layers of the backfill, occurring in association with vivianite. The Roman grave, C51364, was unique 430 among those examined from Hungate in containing fan-like crystals of goethite and an absence of 431 vivianite. The location of the goethite crystals within soil pores around the skull indicate it to be a 432 secondary mineral (Fitzpatrick and Schwertmann, 1982) and the multi-layered fan-like coatings 433 indicate that the soil was very wet (Stoops and Delvigne, 1990). Goethite is formed under changing 434 redox conditions, being produced in intense oxidizing periods (Lindbo et al., 2010). Following oxidative 435 weathering, siderite is converted to goethite, assuming brownish colours (Stoops and Delvigne, 1990).

436 The siderite in the Roman grave formed in the sediments surrounding the skeleton during a phase of 437 reducing conditions, suggested to be associated with the decomposition of the organic matter of the 438 corpse and the coffin under water-logged conditions. The higher frequency of siderite associated with 439 the pelvis can be explained by the greater organic matter content and the consequently more intense 440 reducing sub-environment. Furthermore, the amorphous nature of the cores of the Group 2 441 spherulites may reflect inhibition of crystallisation by the higher amounts of organic matter associated 442 with the torso (cf. Fitzpatrick and Schwertmann, 1982). The orange colour of the siderite is indicative 443 of partial oxidation, presumably resulting as the decay of organic matter slowed reducing oxygen 444 sequestration and/or caused by the excavation, during Anglo-Scandinavian times, of the cesspit 445 located in the proximity of the skull and pelvis and partially cutting the backfill (Fig. 4). The excavation 446 of the cesspit most likely increased water flow through the grave, causing oxygenation. Further 447 evidence for higher fluid flow comes from the amorphous phosphatic coatings around the pore walls in the C3 sample, located in the proximity of the cesspit. During the oxidative phase the sideritic 448 449 spherules from the sediments closest to the cess pit (C3) underwent the most extensive weathering 450 and goethite fan-like crystals were formed within C3 and on the pore walls in the area around the 451 skull. The presence of the Fe-rich corona on the Group 3 spherulites indicates a period of intense 452 anoxia during which the Fe-rich deposits accumulated, consistent with leaching and precipitation of 453 Fe derived from the cess pit following its extended use. The higher Ca contents of the spherulites 454 from the area of the feet can be attributed an origin from their proximity to bone and limited extent of decalcification, the smaller size and higher Fe contents resulting from their rapid sealing and preservation by fine material coatings. Notably, the results from the analysis of the spherulites are entirely consistent with the interpretation from the profiles of the organic signatures and loss of bone from the area of the grave closest to the cess pit.

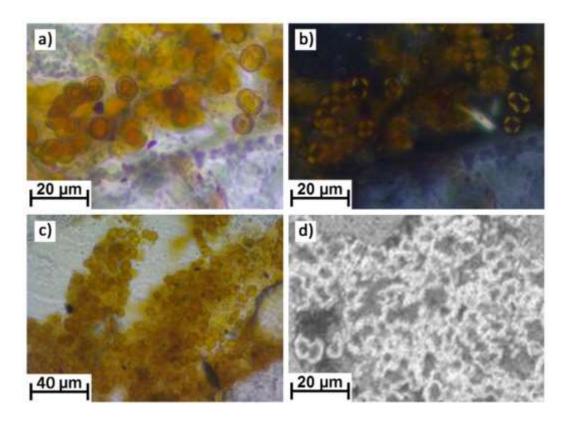


Figure 6. Sideritic spherulites in the Roman age burial C51364: a) and b) from the area of the skull,
showing layered structure in PPL and characteristic Maltese cross extinction in XPL, c) accumulation
in packing voids in the area of the pelvis and d) showing elemental compositional variation with Ferich outer rim in SEM-EDS image.

463 **4 Conclusions**

464 GC-MS and HPLC-MSⁿ analyses revealed distinct spatial variations in lipid profiles within the graves. 465 TAGs, derived from adipose tissue, were observed in the extracts from both graves, being most abundant in samples from around the torso. The TAG acyl chain distributions, dominated by saturated 466 467 components, suggest loss of unsaturated components, either by oxidative cleavage of the unsaturated 468 bonds or by preferential hydrolysis of the unsaturated acyl moieties, rather than reductive 469 transformation into saturated components. The highest levels of DAGs, MAGs, fatty acids and *n*-alkan-470 1-als also represent inputs from degraded body tissues and were observed in extracts from the upper 471 torso and pelvis. *n*-Alkan-1-als, by virtue of their formation from reductive transformation of fatty 472 acids and their survival in the soil, indicate regions of persistent anoxia within the grave. Their

473 preservation is highly localised to regions on the skeleton (e.g. skull, pelvis) where the bone 474 morphology provides a barrier to free movement of water thus can allow anoxic conditions to develop. 475 The higher abundances of microbial markers (branched chain fatty acids and hopanoids) in samples 476 from around the skeleton than in the controls reflects a legacy of increased microbial activity in 477 response to the remains, most likely through their involvement with the degradation of the corpse. 478 In the case of the Roman burial, abundant hopanoids clearly distinguish the gut region of the remains, 479 while their occurrence at the feet indicates migration of some material derived from the gut to the 480 foot of the coffin. The gut region of the Anglo-Scandinavian burial is characterised by elevated levels 481 of the faecal sterol, coprostanol. By contrast, coprostanol is present throughout the Roman grave and 482 instead records the ingress of material from an Anglo-Scandinavian age cess pit adjacent to the burial. 483 The acidic and anoxic nature of the leachate is indicated by the mobilisation of ferrous iron and the 484 formation and persistence of Fe-rich sideritic spherulites. It is clear that leachate from the cess pit has 485 affected the preservation of the skeletal remains, which is notably worse in areas closest to the pit. 486 Decalcification of the bone by acidic fluids from the cess pit has led to release of bone-cholesterol into 487 the soil in those regions, and depleted Ca contents recorded in spherulites from the skull and pelvis. 488 The spherulites record several changes in the redox conditions within the Roman grave, from their 489 initial formation, during a phase of intense reducing conditions associated with decomposition of the 490 corpse, to a phase of partial oxidation as the decay progressed and oxygen demand decreased. This 491 was followed by a second, more intensive phase of oxidation associated with the construction of the 492 cess pit and increased water flow through the grave, indicated by goethite formation and phosphatic 493 coatings on pore walls in the controls. An intense period of anoxia then followed with the extended 494 usage of the cess pit, marked by deposition on the spherulites of iron rich layers.

495 The results exemplified by analysis of the two graves demonstrate that a valuable unexploited archive 496 of information pertinent to the taphonomic interpretation of human burials is recorded in the burial 497 matrix. The results demonstrate the preservation of signatures relating to the body tissues and gut 498 contents which have the potential to inform archaeological investigation of diet (body tissues) and 499 last meals (gut contents) both within particular communities and between different communities. 500 Furthermore, the recognition that such organic signatures can and do survive should provide a 501 stimulus to further research. Notably, the potential for survival of signatures from grave offerings 502 provides scope for gaining further insights into their nature and use. Understanding of the factors 503 that influence the survival/destruction of organic signatures in graves could also reveal characteristics 504 that relate to particular styles of burial. The methods described here could also have potential in the 505 identification of temporary or disinterred graves including in determining the anatomical orientation 506 from the chemical signatures preserved in the soil, though further research is necessary.

507 Acknowledgements

508 The research forms part of the InterArChive project and received funding from the European Research

509 Council under the European Community's Seventh Framework Programme (FP7/2007-2013) / ERC

- 510 grant agreement n° 230193. We would like to express our sincere thanks to York Archaeological Trust
- 511 for allowing us to collect samples. Karl Heaton is thanked for acquiring the GC-MS data.

512 References

- 513 Barfield, L., 1994. The Iceman reviewed. Antiquity 68, 10–26.
- Bethell, P.H., Goad, L.J., Evershed, R.P., Ottaway, J., 1994. The study of molecular markers of human
 activity the use of coprostanol in the soil as an indicator of human fecal material. J. Archaeol.
 Sci. 21, 619–632.
- 517 Brothwell, D.R., 1981. Digging up Bones: the Excavation, Treatment and Study of Human Skeletal 518 Remains, 3rd ed. British Museum (Natural History) and Oxford University Press.
- Browne, G.H., Kingston, D.M., 1993. Early diagenetic spherulitic siderites from Pennsylvanian
 palaeosols in the Boss Point Formation, Maritime Canada. Sedimentology 40, 467–474.
 https://doi.org/10.1111/j.1365-3091.1993.tb01346.x
- 522 Buckley, S.A., Evershed, R.P., 2001. Organic chemistry of embalming agents in Pharaonic and Graeco-523 Roman mummies. Nature 413, 837–841.
- Bull, I.D., Berstan, R., Vass, A., Evershed, R.P., 2009. Identification of a disinterred grave by molecular
 and stable isotope analysis. Sci. Justice 49, 142–149.
 https://doi.org/10.1016/j.scijus.2009.01.016
- Bullock, P., Fedoroff, N., Jongerius, A., Stoops, G., Tursina, T., 1985. Handbook for Soil Thin Section
 Description. Waine Research Wolverhampton.
- Burns, A., Pickering, M.D., Green, K.A., Pinder, A.P., Gestsdóttir, H., Usai, M.-R., Brothwell, D.R., Keely,
 B.J., 2017. Micromorphological and chemical investigation of late-Viking age grave fills at
 Hofstaðir, Iceland. Geoderma 306, 183–194. https://doi.org/10.1016/j.geoderma.2017.06.021
- Can, I., Javan, G.T., Pozhitkov, A.E., Noble, P.A., 2014. Distinctive thanatomicrobiome signatures found
 in the blood and internal organs of humans. J. Microbiol. Methods 106, 1–7.
 https://doi.org/10.1016/j.mimet.2014.07.026
- 535 Courty, M.A., Goldberg, P., Macphail, R.I., 1989. Soils and Micromorphology in Archaeology.
 536 Cambridge University Press Cambridge.
- 537 Cross, A.F., Schlesinger, W.H., 1995. A literature review and evaluation of the. Hedley fractionation:

- 538 Applications to the biogeochemical cycle of soil phosphorus in natural ecosystems. Geoderma 539 64, 197–214. https://doi.org/10.1016/0016-7061(94)00023-4
- 540 Driese, S.G., Ludvigson, G.A., Roberts, J.A., Fowle, D.A., Gonzalez, L.A., Smith, J.J., Vulava, V.M., McKay,
- 541 L.D., 2010. Micromorphology and Stable-Isotope Geochemistry of Historical Pedogenic Siderite
- 542 Formed in PAH-Contaminated Alluvial Clay Soils, Tennessee, U.S.A. J. Sediment. Res. 80, 943–
- 543 954. https://doi.org/10.2110/jsr.2010.087
- Eglinton, G., Hamilton, R.J., 1967. Leaf epicuticular waxes. Science (80-.). 156, 1322–1335.
 https://doi.org/10.1126/science.156.3780.1322
- Eigenbrode, J.L., Freeman, K.H., Summons, R.E., 2008. Methylhopane biomarker hydrocarbons in
 Hamersley Province sediments provide evidence for Neoarchean aerobiosis. Earth Planet. Sci.
 Lett. 273, 323–331. https://doi.org/10.1016/j.epsl.2008.06.037
- 549Evershed, R.P., 2008. Organic residue analysis in archaeology: the archaeological biomarker550revolution. Archaeometry 50, 895–924. https://doi.org/10.1111/j.1475-4754.2008.00446.x
- Evershed, R.P., Connolly, R.C., 1988. Lipid preservation in Lindow man. Naturwissenschaften 75, 143–
 145. https://doi.org/10.1007/bf00405306
- Evershed, R.P., Turnerwalker, G., Hedges, R.E.M., Tuross, N., Leyden, A., 1995. Preliminary-results for
 the analysis of lipids in ancient bone. J. Archaeol. Sci. 22, 277–290.
- 555 Fitzpatrick, E.A., 1993. Soil microscopy and micromorphology. John Wiley & Sons Ltd, West Sussex.
- Fitzpatrick, R.W., Schwertmann, U., 1982. Al-Substituted goethite--an indicator of pedogenic and
 other weathering environments in South Africa. Geoderma 27, 335–347.
- Forbes, S.L., Stuart, B.H., Dent, B.B., 2002. The identification of adipocere in grave soils. Forensic Sci.
 Int. 127, 225–230. https://doi.org/10.1016/S0379-0738(02)00127-5
- Garnier, N., Rolando, C., Høtje, J.M., Tokarski, C., 2009. Analysis of archaeological triacylglycerols by
 high resolution nanoESI, FT-ICR MS and IRMPD MS/MS: Application to 5th century BC-4th
 century AD oil lamps from Olbia (Ukraine). Int. J. Mass Spectrom. 284, 47–56.
 https://doi.org/10.1016/j.ijms.2009.03.003
- Gebhardt, A., Langohr, R., 1999. Micromorphological Study of Construction Materials and Living Floors
 in the Medieval Motte of Werken (West Flanders, Belgium). Geoarchaeology An Int. J. 14, 595–
 620. https://doi.org/10.1002/(SICI)1520-6548(199910)14:7<595::AID-GEA1>3.0.CO;2-Q
- 567 Ghislandi, S., 2016. Graves under the microscope: micromorphological study of sediments in 568 archaeological burials. University of York.

- Gilbertson, J.R., Ferrell, W.J., Gelman, R.A., 1967. Isolation and analysis of free fatty aldehydes from
 rat, dog ,and bovine heart muscle. J. Lipid Res. 8, 38–45.
- 571 Goldberg, P., Macphail, R.I., 2003. Short contribution: Strategies and techniques in collecting 572 micromorphology samples. Geoarchaeology 18, 571–578. https://doi.org/10.1002/gea.10079
- 573 Grimalt, J.O., Fernandez, P., Bayona, J.M., Albaiges, J., 1990. Assessment of fecal sterols and ketones
- as indicators of urban sewage inputs to coastal waters. Environ. Sci. Technol. 24, 357–363.
 https://doi.org/10.1021/es00073a011
- 576 Gulaçar, F.O., Susini, A., Klohn, M., 1990. Preservation and postmortem transformations of lipids in 577 samples from a 4000-year-old Nubian mummy. J. Archaeol. Sci. 17, 691–705.
- Hadian, M., Good, I., Pollard, A.M., Zhang, X., Laursen, R., 2012. Textiles from Douzlakh Salt Mine at
 Chehr Abad, Iran: a technical and contextual study of Late pre-Islamic Iranian textiles. Int. J.
 Humanit. Islam. Repub. Iran 19, 152–173.
- Hasan, H., 2010. Development of an LC-MS/MS Method for the Analysis of Triacylglycerols from Meat
 and Application in the Discrimination of Cooked Meat Products. Ph.D. Thesis, Univ. York. PhD
 thesis, University of York UK.
- Hita, C., Parlanti, E., Jambu, P., Joffre, J., Amblès, A., 1996. Triglyceride degradation in soil. Org.
 Geochem. 25, 19–28. https://doi.org/10.1016/S0146-6380(96)00107-6
- Hodson, L., Skeaff, C.M., Fielding, B.A., 2008. Fatty acid composition of adipose tissue and blood in
 humans and its use as a biomarker of dietary intake. Prog. Lipid Res. 47, 348–380.
 https://doi.org/10.1016/j.plipres.2008.03.003
- Jensen, R.G., Dejong, F.A., Lambertdavis, L.G., Hamosh, M., 1994. Fatty-acid and positional selectivities
 of gastric lipase from premature human infants in-vitro studies. Lipids 29, 433–435.
 https://doi.org/10.1007/bf02537313
- Jim, S., Ambrose, S.H., Evershed, R.P., 2004. Stable carbon isotopic evidence for differences in the
 dietary origin of bone cholesterol, collagen and apatite: implications for their use in
 palaeodietary reconstruction. Geochim. Cosmochim. Acta 68, 61–72.
 https://doi.org/10.1016/s0016-7037(03)00216-3
- Kimpe, K., Jacobs, P.A., Waelkens, M., 2001. Analysis of oil used in late Roman oil lamps with different
 mass spectrometric techniques revealed the presence of predominantly olive oil together with
 traces of animal fat. J. Chromatogr. A 937, 87–95. https://doi.org/10.1016/S00219673(01)01304-8

- Leeming, R., Ball, A., Ashbolt, N., Nichols, P., 1996. Using faecal sterols from humans and animals to
 distinguish faecal pollution in receiving waters. Water Res. 30, 2893–2900.
- Lemos, V.P., Da Costa, M.L., Lemos, R.L., De Faria, M.S.G., 2007. Vivianite and siderite in lateritic iron
 crust: An example of bioreduction. Quim. Nova 30, 36–40. https://doi.org/10.1590/S010040422007000100008
- Lindbo, D.L., Stolts, M.H., Vepraskas, M.L., 2010. Redoximorphic Features, in: Stoops, G., Marcelino,
 V., Mees, F. (Eds.), Interpretation of Micromorphological Features of Soils and Regoliths. Elsevier,
 UK.
- Ludvigson, G.A., González, L.A., Fowle, D.A., Roberts, J.A., Driese, S.G., Villarreal, M.A., Smith, J.J.,
 Suarez, M.B., 2013. Paleoclimatic Applications and Modern Process Studies of Pedogenic
 Siderite. New Front. Paleopedol. Terr. Paleoclimatology 79–87.
 https://doi.org/10.2110/sepmsp.104.01
- Mayer, B.X., Reiter, C., Bereuter, T.L., 1997. Investigation of the triacylglycerol composition of
 Iceman's mummified tissue by high-temperature gas chromatography. J. Chromatogr. B 692, 1–
 6.
- 615 Miedema, R., Pape, T., Van der Waal, G.J., 1974. A method to impregnate wet soil samples, producing
 616 high-quality thin sections. Netherlands J. Agric. Sci. 22, 37–39.
- Milek, K.B., 2012. Floor formation processes and the interpretation of site activity areas: An
 ethnoarchaeological study of turf buildings at Thverá, northeast Iceland. J. Anthropol. Archaeol.
 31, 119–137.
- Mirabaud, S., Rolando, C., Regert, M., 2007. Molecular criteria for discriminating adipose fat and milk
 from different species by nanoESI MS and MS/MS of their triacylglycerols: Application to
 archaeological remains. Anal. Chem. 79, 6182–6192. https://doi.org/10.1021/ac070594p
- Nriagu, J., 1972. Stability of vivianite and ion-pair formation in the system Fe₃(PO₄)₂-H₃PO₄H₃PO₄-H₂O.
 Geochim. Cosmochim. Acta 36, 459–470.
- O'Connor, S., Ali, E., Al-Sabah, S., Anwar, D., Bergström, E., Brown, K.A., Buckberry, J., Buckley, S.,
 Collins, M., Denton, J., Dorling, K.M., Dowle, A., Duffey, P., Edwards, H.G.M., Faria, E.C., Gardner,
 P., Gledhill, A., Heaton, K., Heron, C., Janaway, R., Keely, B.J., King, D., Masinton, A., Penkman,
 K., Petzold, A., Pickering, M.D., Rumsby, M., Schutkowski, H., Shackleton, K.A., Thomas, J.,
 Thomas-Oates, J., Usai, M.-R., Wilson, A.S., O'Connor, T., 2011. Exceptional preservation of a
 prehistoric human brain from Heslington, Yorkshire, UK. J. Archaeol. Sci. 38, 1641–1654.
 https://doi.org/10.1016/j.jas.2011.02.030

- Pickering, M.D., Lang, C., Usai, M.-R., Keely, B.J., Brothwell, D.R., 2014. Organic residue analysis of
 soils, in: Loe, L., Boyle, A., Webb, H., Score, D. (Eds.), "Given to the Ground": A Viking Age Mass
 Grave on Ridgeway Hill, Weymouth. Dorset Natural History and Archaeological Society
 Monograph Series Vol 23, Oxbow Books Oxford, pp. 237–245.
- Pye, K., Dickson, J.A.D., Schiavon, N., Coleman, M.L., Cox, M., 1990. Formation of Siderite Mg-Calcite
 Iron Sulfide Concretions in Intertidal Marsh and Sandflat Sediments, North-Norfolk, England.
 Sedimentology 325–343.
- Quirk, M.M., Wardroper, A.M.K., Wheatley, R.E., Maxwell, J.R., 1984. Extended hopanoids in peat
 environments. Chem. Geol. 42, 25–43. https://doi.org/10.1016/0009-2541(84)90003-2
- Raclot, T., Groscolas, R., 1993. Differential mobilization of white adipose-tissue fatty-acids according
 to chain-length, unsaturation, and positional isomerism. J. Lipid Res. 34, 1515–1526.
- Rapport, M.M., Lerner, B., 1959. The structure of plasmalogens IV. Lipids in normal and neoplastic
 tissues of man and in normal tissues of rabbit and rat. Biochim. Biophys. Acta 33, 319–325.
 https://doi.org/10.1016/0006-3002(59)90119-2
- Šoberl, L., Gasparic, A.Z., Budja, M., Evershed, R.P., 2008. Early herding practices revealed through
 organic residue analysis of pottery from the early Neolithic rock shelter of Mala Triglavca ,
 Slovenia. Doc. Praehist. 15, 253–260. https://doi.org/10.4312/dp.35.19
- Stead, I.M., Bourke, J.B., Brothwell, D.R., 1986. Lindow Man: The Body in the Bog. British Museum
 Publications, London.
- Stoops, G., 2003. Guidelines for Analysis and Description of Soil and Regolith Thin-Sections. SoilScience Society of America Madison.
- Stoops, G., Delvigne, J., 1990. Morphology of Mineral Weathering and Neoformation. II
 Neoformations. Dev. Soil Sci. 19, 483–492.
- Stott, A.W., Evershed, R.P., Jim, S., Jones, V., Rogers, J.M., Tuross, N., Ambrose, S., 1999. Cholesterol
 as a new source of palaeodietary information: experimental approaches and archaeological
 applications. J. Archaeol. Sci. 26, 705–716. https://doi.org/10.1006/jasc.1998.0386
- Winkler, A., Haumaier, L., Zech, W., 2001. Variation in hopanoid composition and abundance in forest
 soils during litter decomposition and humification. Org. Geochem. 32, 1375–1385.
 https://doi.org/10.1016/s0146-6380(01)00115-2
- Wittenberg, J.B., Korey, S.R., Swenson, F.H., 1956. Determination of higher fatty aldehydes in tissues.
 J. Biol. Chem. 219, 39–47.

663